

# Using Observations of HNO<sub>3</sub> and N<sub>2</sub>O to Quantify HCl and Ozone Sensitivity to Variability of the Stratospheric Circulation

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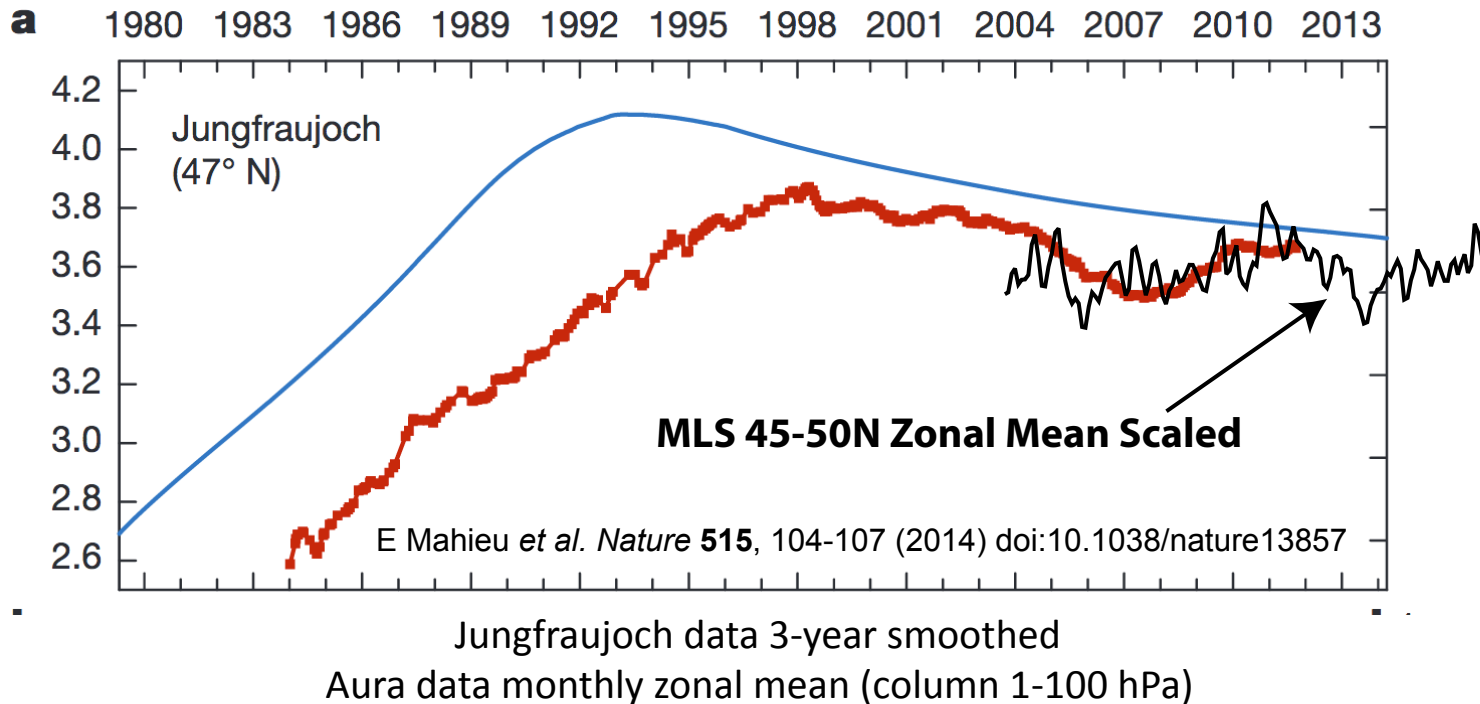


# Why do we care about interannual variability in stratospheric dynamics?

- **Interannual variability masks detection of trends**
  - a) Masks recovery of chlorine as measured by HCl column amounts
  - b) complicates detection of expected upward trend in total ozone or lower stratospheric ozone due to chlorine change
- **On longer time scales, models predict speedup of BDC**
  - a) No clear confirmation of these predictions by measurements
  - b) Interannual variability of dynamics masks slow predicted change



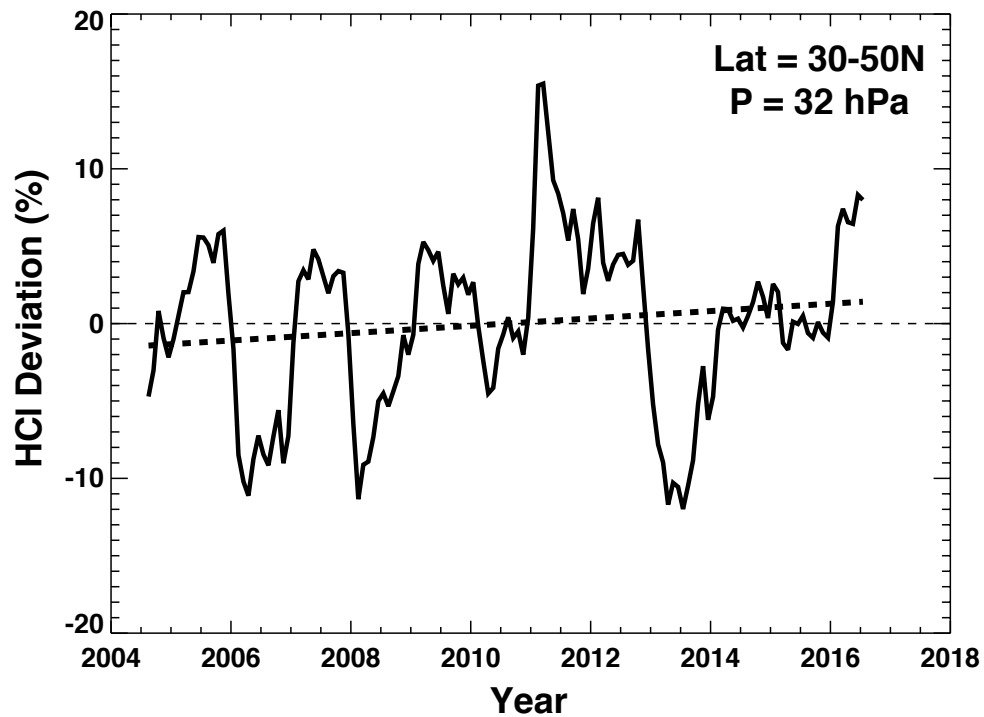
# Motivation: HCl column variation at northern



Can we find surrogate for dynamical influence on HCl variation?



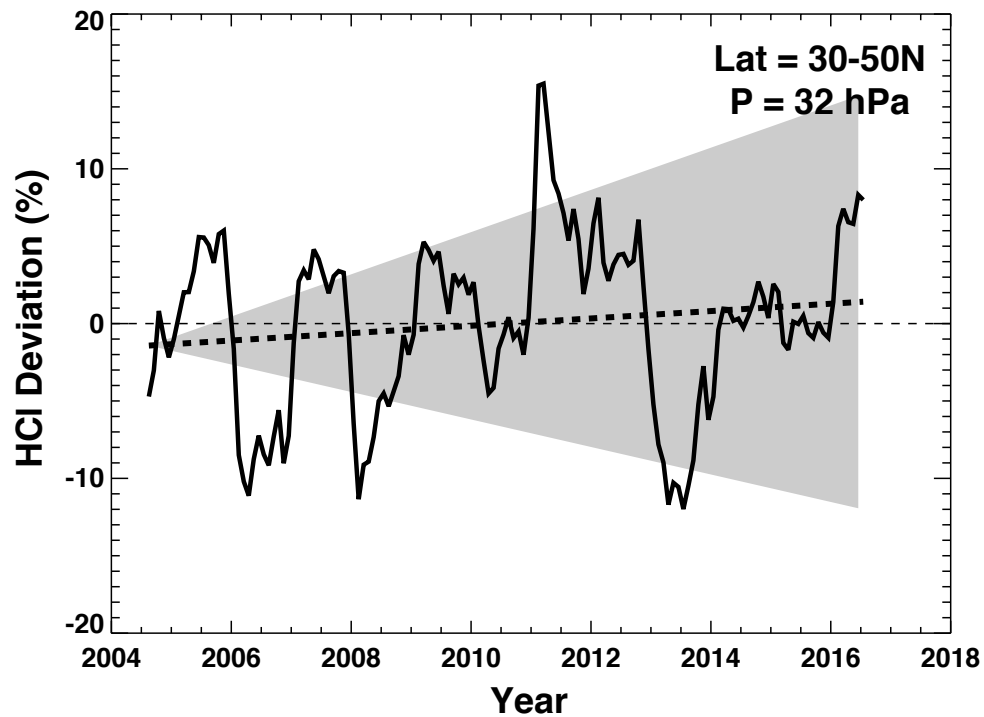
# Deseasonalized HCl Anomalies at 32 hPa for 30-50N Latitude Band



Small Positive Trend



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Small Positive Trend

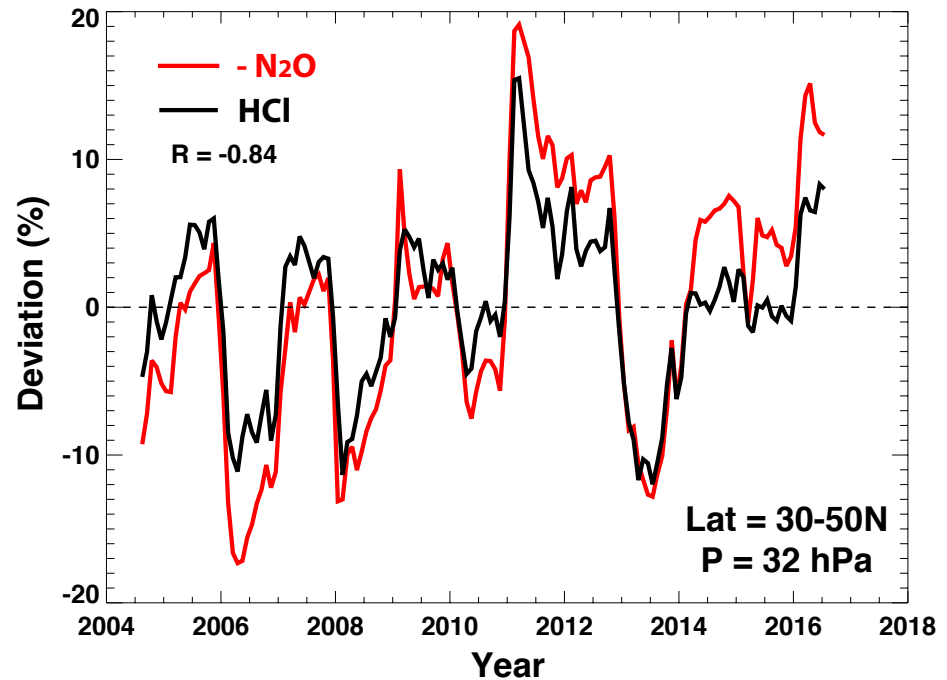
No Statistical Significance



## Consider altitude profiles of N<sub>2</sub>O and HCl from Aura MLS

- N<sub>2</sub>O and HCl both respond to dynamical changes through their spatial gradients
- They are anti-correlated at a given latitude and pressure level

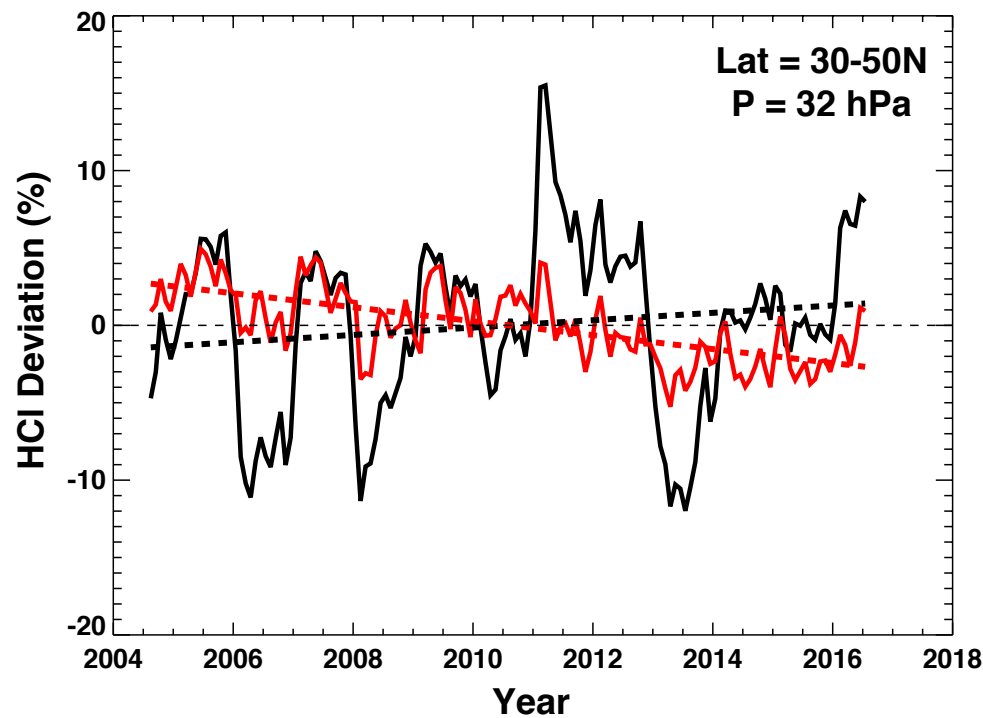
Suggests that N<sub>2</sub>O variations could be used to model/remove variability in HCl observations to reveal trend



*Monthly mean anomalies of N<sub>2</sub>O and HCl from Aura MLS measurements at 32 hPa averaged between latitudes of 30 to 50N.*

Douglass, Strahan, and Stolarski, QOS Edinburgh, 5 Sept. 2016

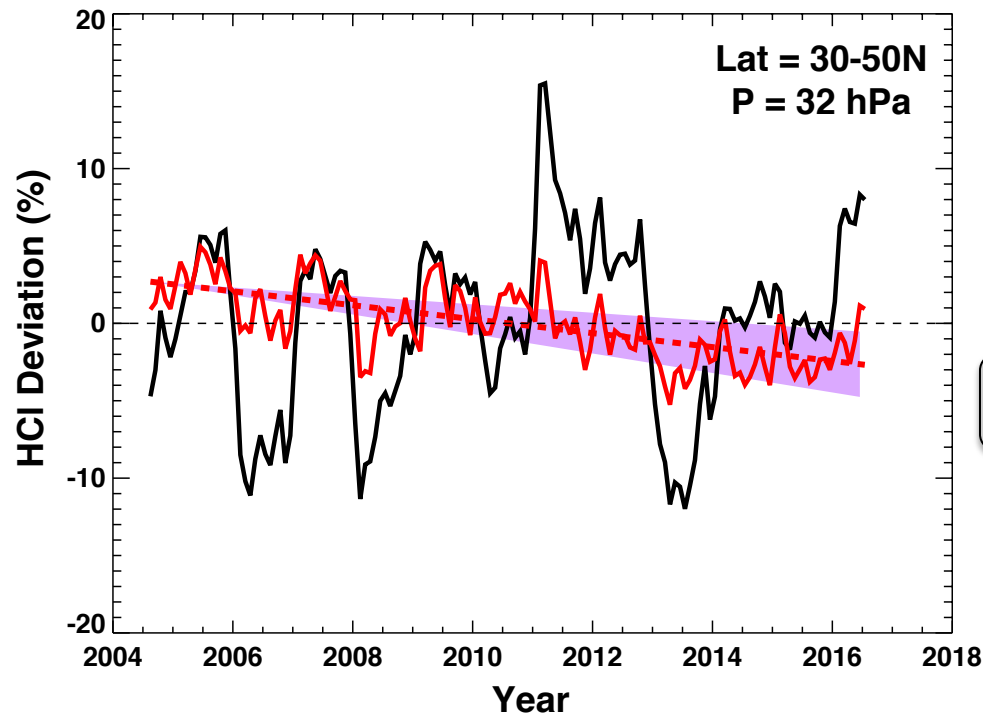
# HCl Anomalies with seasonal cycle and N<sub>2</sub>O co-variation removed



Negative Trend



# HCl Anomalies with seasonal cycle and N<sub>2</sub>O co-variation removed



Negative Trend

Statistically Significant at  $2\sigma$





# Fit linear trend to HCl time series at each altitude of reported MLS measurements

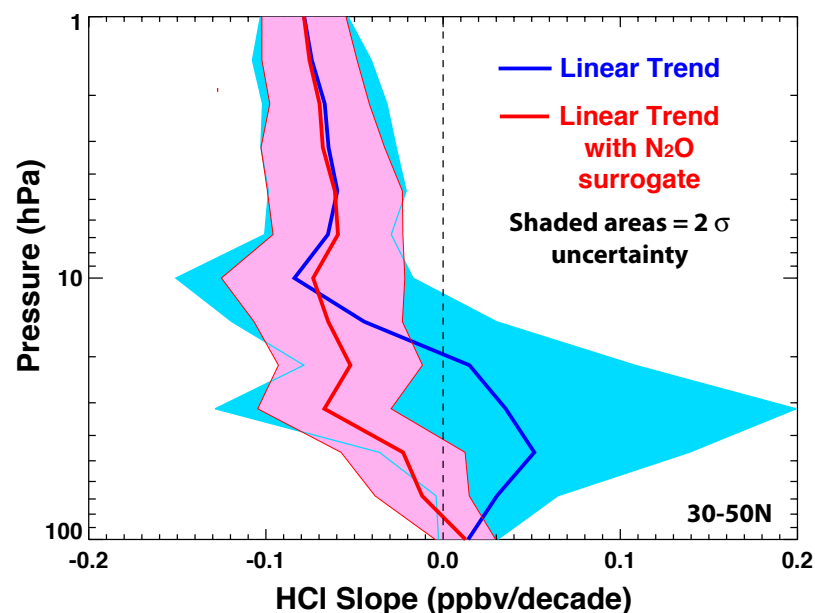
- First use simple linear trend model

$$\text{HCl} = \mu + \alpha \cdot \text{trend} + \beta \cdot \text{seasonal} + \varepsilon$$

- Then add term for N<sub>2</sub>O anomalies

$$\text{HCl} = \mu + \alpha \cdot \text{trend} + \beta \cdot \text{seasonal} + \gamma \cdot \text{N}_2\text{O} + \varepsilon$$

Trend changes sign with smaller uncertainty



*HCl slope from August 2004 through August 2016 from MLS data between 30 and 50 N latitudes. Shaded areas are 2σ uncertainty estimates for trend.*

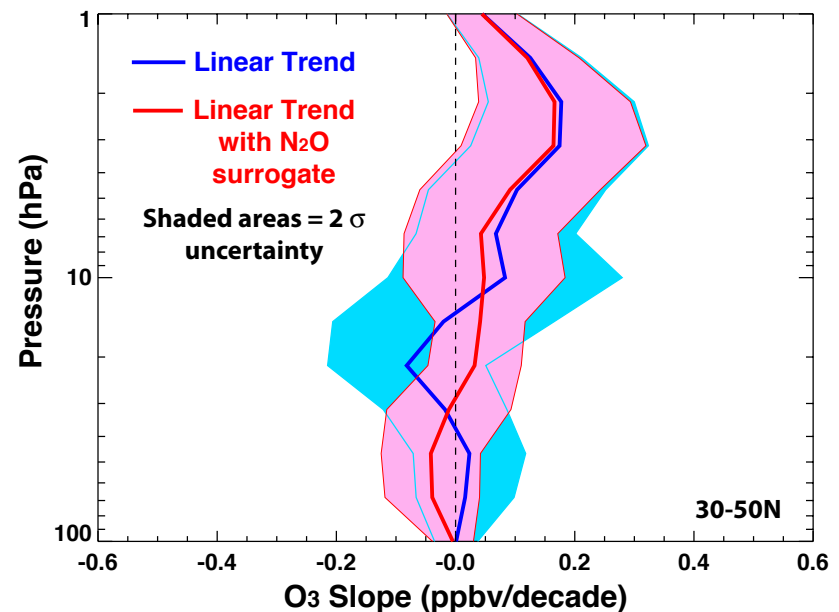
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## What about O<sub>3</sub>?

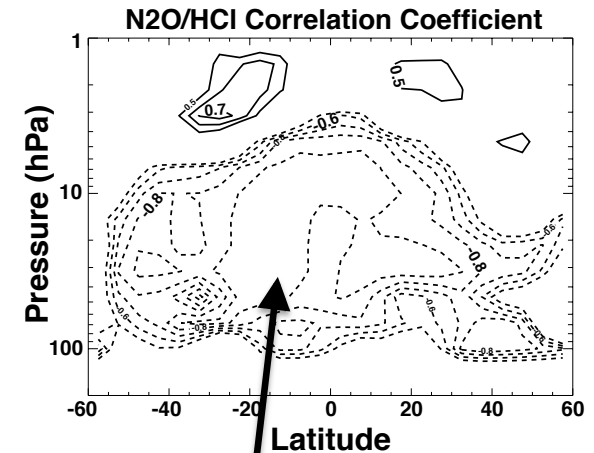
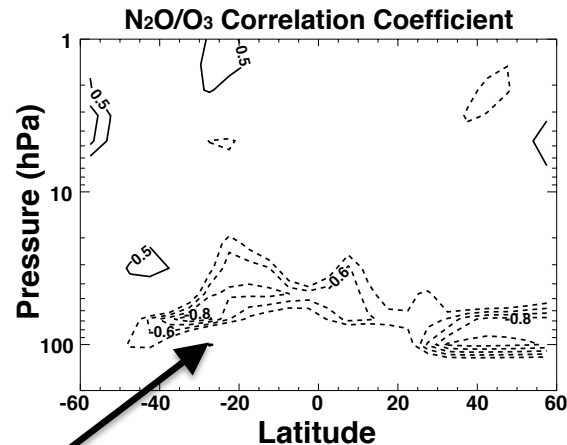
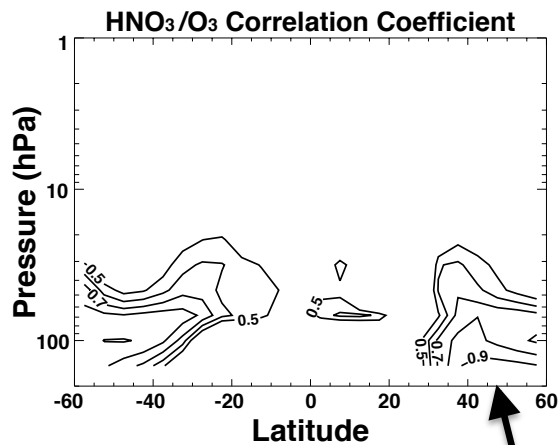
- Using N<sub>2</sub>O as fitting term reduces uncertainty in O<sub>3</sub> trend
- Using HNO<sub>3</sub> yields similar results (not shown)
- Calculated trend becomes positive in middle stratosphere as expected, but results are not significant

Accounting for dynamical variability in O<sub>3</sub> trends will be more difficult



*O<sub>3</sub> slope from August 2004 through August 2016 from MLS data between 30 and 50 N latitudes. Shaded areas indicate 2σ uncertainty estimates of the trend.*

# Usefulness of dynamical tracer depends on correlations that are determined by gradients



O<sub>3</sub> and N<sub>2</sub>O or HNO<sub>3</sub> are strongly correlated only in the lower stratosphere

HCl and N<sub>2</sub>O strongly anti-correlate over most of stratosphere

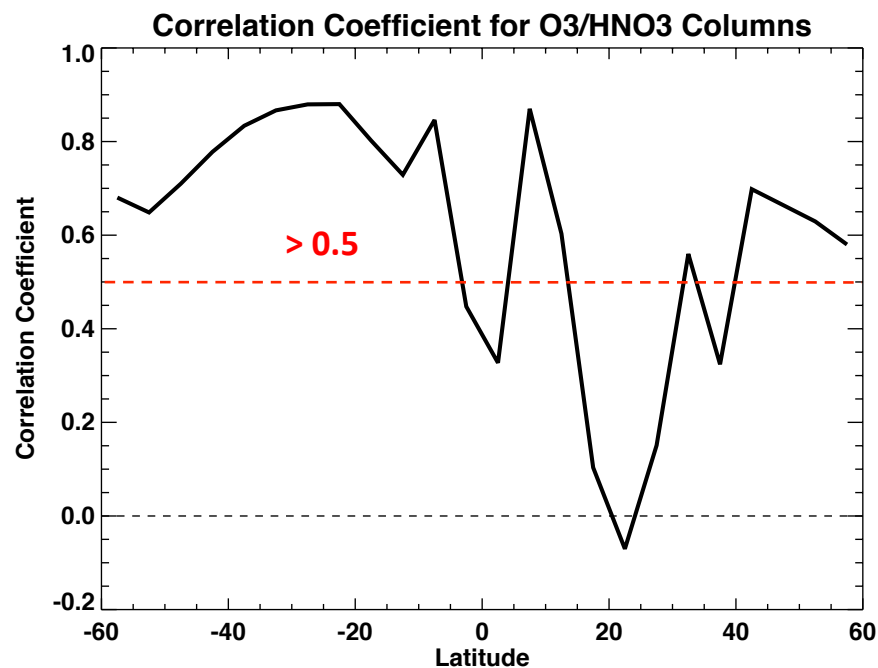
Only correlations < -0.5 and > 0.5 shown



# Can we do better using column $\text{HNO}_3$ with column $\text{O}_3$ ?

- Correlation is  $> 0.5$  over entire SH and between 40-60N
- Warrants further examination

Would be very useful as we could extend study back in time using column measurements from NDACC stations



# Conclusions

- Dynamical variability introduces uncertainty into trend analysis of chemical constituents such as O<sub>3</sub> and HCl
- Many studies have used dynamical surrogates such as QBO, ENSO, AMO in trend models to try to remove (explain) this variance
- We propose using constituent correlations to accurately model the “whole dynamical” impact on species variability
- We have shown important example of removing variability in HCl measurements from Aura MLS by using measurements of N<sub>2</sub>O

